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Arnaud Mialon, Jean-Pierre Wigneron, Patricia De Rosnay, Maria-José Escorihuela, Y.H. Kerr. Evaluating the L-MEB Model From Long-Term Microwave Measurements Over a Rough Field, SMOSREX 2006. IEEE Transactions on Geoscience and Remote Sensing, Institute of Electrical and Electronics Engineers, 2012, vol.50 (no.5), pp.1458-1467. <10.1109/TGRS.2011.2178421>. <hal-00690897>

HAL Id: hal-00690897

<https://hal.archives-ouvertes.fr/hal-00690897>

Submitted on 25 Apr 2012

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Evaluating the L-MEB model from long term microwave measurements over a rough field, SMOSREX 2006

Arnaud Mialon, Jean-Pierre Wigneron, *Senior Member 03, IEEE*, Patricia de Rosnay, Maria Jose Escorihuela, Yann H. Kerr, *Member 88, Senior Member 01, IEEE*

Abstract—The present study analyses the effects of the roughness on the surface emission at L-band, based on observations acquired during a long term experiment. At the SMOSREX (Surface Monitoring Of the Soil Reservoir EXperiment) site near Toulouse, France, a bare soil was ploughed and monitored over more than a year by means of a L-band radiometer profile soil moisture and temperature sensors as well, as a local weather station, accompanied by 12 roughness campaigns. The aim of this study is (1) to present this unique database, and (2) to use this dataset to investigate the semi-empirical parameters for the roughness in L-MEB (L-Band Microwave Emission of the Biosphere), that is the forward model used, in the SMOS (Soil Moisture and Ocean Salinity) soil moisture retrieval algorithm. In particular, we studied the link between these semi empirical parameters and the soil roughness characteristics expressed in terms of standard deviation of surface height (σ) and the correlation length (LC). The dataset verifies that roughness effects decrease the sensitivity of surface emission to soil moisture, an effect which is most pronounced at high incidence angles and soil moisture and at horizontal polarization. Contradictory to previous studies, the semi-empirical parameter Q_r was not found to be equal to 0 for rough conditions. A linear relationship between the semi-empirical parameters N and σ was established, while N_H and N_V appeared to be lower for a rough ($N_H \sim 0.59$ and $N_V \sim -0.3$) than for a quasi-smooth surface. This study reveals the complexity of roughness effects and demonstrates the great value of a sound long-term dataset of rough L-band surface emissions to improve our understanding on the matter.

Index Terms—SMOS, Roughness, Passive Microwave, L-band, L-MEB model.

I. INTRODUCTION

SOIL moisture is a key parameter controlling air-land interface exchanges. Although very

important in many applications (climate models, agriculture, water resources management), it is difficult to monitor this variable at a global scale. The SMOS (Soil Moisture and Ocean Salinity) satellite mission [1; 2], successfully launched in November 2009, is the first mission to deliver global surface soil moisture fields at a high temporal resolution of 3 days. The retrieval scheme to derive soil moisture [3] is based on multi-angular passive microwave brightness temperatures ($f=1.4$ GHz) as measured by the instrument [4] and on surface emission models at L-band (L-MEB, L-band Microwave Emission of the Biosphere [5; 6; 7]).

Land surface emission at this wavelength is mainly controlled by soil moisture but important issues are still to be tackled [8] such as roughness, which is the focus of this paper. Roughness influence on surface emission is complex as it implies 3-D geometric soil surface features as well as soil moisture heterogeneity, in particular between peaks and hollows. Its major effect is to decrease the sensitivity of L-band brightness temperatures to soil moisture [9; 10]. Shi et al. [11] found by the use of an Integral Equation Model (IEM) that roughness influence is more significant at high incidence angles and high soil moisture content as well as a function of polarization. They noted an increase in emissivity with roughness at the horizontal polarization at low incidence angles. For dry soil, the emissivity of the vertical polarization (typically higher than ~ 0.8) shows a decrease compared with that of a flat surface, whereas for wet soil (emissivity lower than ~ 0.8) an increase is observed.

Using complex models as the IEM approach to compute the surface emissivity is not possible in the SMOS soil moisture algorithm as it needs many inputs and its computation is time demanding. Instead the SMOS level2 retrieval algorithm [3] uses semi-empirical approaches [7; 8] to compute the emission of the surface. The correction for a rough surface [9; 10; 12; 13]

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is based on empirical parameters (H_r , N_V , N_H , Q_r) that have to be calibrated with in-situ data reflecting local surface characteristics (soil texture level of roughness). Most recent studies on L-band emission [14; 15; 16] have retrieved these parameters to best fit the observations, but more investigations are needed on roughness to relate the soil L-MEB parameters to the surface roughness characteristics.

The roughness analyses conducted so far have all been either restricted to short investigation periods [17; 13] or to almost flat surface conditions [18] only. These have motivated the present study which for the first time focuses on a rough soil observed over a long time period at the SMOSREX (Surface Monitoring Of the Soil Reservoir EXperiment) site in 2006/07. A bare soil was ploughed creating a very rough surface and its roughness evolved over more than a year naturally due to climatic events (rainfalls, wind).

The aim of this study is twofold. First, this unique database (referred to as SMOSREX-2006) is presented and the L-band observations over the rough surface covering a wide range of soil moisture conditions (from very wet to very dry) are analysed over a long period of time (14 months). Second, the SMOSREX-2006 is used to evaluate the roughness parameters of the semi empirical model used in the L-MEB model. Q_r , H_r and N_V (p for the polarization horizontal or vertical) are retrieved in this evaluation and compared with the surface roughness characteristics.

II. MATERIAL

A. Database and experimental site

In preparation of the SMOS mission, the experimental site of SMOSREX (Surface Monitoring Of the Soil Reservoir EXperiment [19]) has been set up near Toulouse in the Southwest of France. Operating since 2003, the database has been used to improve the models implemented in the SMOS soil moisture retrieval [3; 20; 18].

It is equipped with the LEWIS (L-band radiometer for Estimating Water in Soil) radiometer [21], which has been continuously monitoring the emission of the surface. The instrument, placed on a 15m high tower can monitor two fields one with grasscover and a bare soil. It acquires brightness temperatures at vertical and horizontal polarizations (commonly referred to as V and H) at the same frequency as SMOS, i.e. 1.4 GHz, at several incidence angles (i.e. 20, 30, 40, 50 and 60°) every 3 hours (i.e. 2h30, 5h30, 8h30, 11h30,

14h30, 17h30, 20h30, 23h30 UTC).

Additionally, ground measurements are available. Soil texture was analysed and the bare soil was found to be 17% clay, 36% sand and 47% silt [19]. The SMOSREX site is equipped with a weather station, which has been monitoring meteorological data (air temperature, pressure, precipitation, wind) and soil moisture and temperature profiles are measured on each field every 30 minutes. Temperatures measured at different depths, i.e. at 1cm, 5cm, 20cm, 50cm and 90cm with one probe per depth, at the same location as the soil moisture probes, are used to compute the soil temperature. Surface soil moisture is obtained by averaging data from 5 probes placed at the surface (top 0-6 cm layer) on the bare soil field. Soil moisture probes are calibrated from gravimetric measurements [22], from which soil density is estimated.

It is important to note that obtaining an accurate estimation of soil moisture is difficult and can be slightly different from what contributes to the brightness temperatures measured by the radiometer. Due to surface heterogeneity, some differences can occur between the surface covered by the probes ($\sim 4m^2$) and LEWIS field of view that covers a wider surface [19]. Moreover, peaks and hollows imply strong heterogeneity in the surface soil moisture conditions, as soil water content is generally higher in hollows than on peaks. Finally, soil moisture probes measure the dielectric constant over the 0-6 cm top soil layer, whereas the surface emission in L-band is expected to be correlated to the soil moisture of the top 2-3 cm soil layer [23].

B. Roughness measurements

The roughness experiment took place on the bare soil field. On January 13th, 2006 the field was ploughed in a deep manner to ensure a distinct row structure parallel to LEWIS plane of incidence. Thereafter, surface roughness changed naturally over time in response to climatic events, mainly rainfalls and wind.

Surface roughness is measured by means of a two meter long needle board with 201 needles at 1 cm spacing. The needles move freely in the vertical direction and were allowed to fall till they touched the surface reproducing surface variations. Twelve measurement campaigns were conducted over the following 14 months (see Table I), each consisting in the acquisition of several roughness profiles (up to 6), in both directions, i.e. parallel and perpendicular to the plane of incidence of

the LEWIS instrument. Pictures of each vertical profiles were taken with a digital camera to obtain the corresponding numerical profiles of the height variation. These were then used to derive two statistical parameters describing the surface, the standard deviation of heights σ and the correlation length LC [24]. The daily σ are obtained by averaging the variance, i.e. σ^2 , of the different samples acquired in both directions. LC was derived from the auto-correlation function $C(x)$, Eq. 1 [24; 25], which measures the correlation between two heights separated by a distance x :

$$C(x) = \frac{\sum_{i=1}^{N+1-j} z_i \cdot z_{j+1-i}}{\sum_{i=1}^N z_i^2} \quad (1)$$

where $z(i)$ is the height of the needles; j integer ≥ 1 ; the spatial displacement $x=(j-1) \cdot \delta x$; δx being the distance between 2 needles, i.e. 1 cm; N the number of needles $N=201$. The LC corresponds to the distance x where the correlation function (Eq. 1) has decreased to $1/e$, i.e. beyond which two heights are no longer statistically correlated [24]. The auto-correlation function is commonly approximated by the function $C(x)=\exp\left(\frac{-|x|^n}{LC^n}\right)$ where $n=1$ for the Exponential model or $n=2$ for the Gaussian model [26; 25; 27]. For each day of measurement LC is simply the average of the different profiles, mixing both directions. For example, a flat surface is characterized by a low σ and a high LC.

Data acquired before this campaign, i.e. in February and April 2004 [18] and in January 13th just before ploughing the soil, are also used as they provide additional information concerning a quasi-smooth surface. Roughness was also measured in 2010 so that the soil roughness temporal variation could be estimated at interannual scale.

III. METHODOLOGY

A. Observations

The first part of our study is dedicated to surface emission at L-band as observed by the LEWIS radiometer. All cases such as freezing, snow storm on January 28-30 2006) that may introduce artefacts are excluded from the dataset. It is more pertinent to study surface emissivity than brightness temperature as the latter is also influenced by the soil temperature. The emissivity ϵ of a bare soil is obtained from the measured brightness temperatures by removing surface temperature and the sky contributions by applying the following $\epsilon_p = (TB_p - TB_{sky}) / (T_{eff} - TB_{sky})$, where the subscript p stands for the polarization (H or V), and T_{eff}

the effective soil temperature [28] as computed from measured temperatures at all depths based on [19; 29]. The sky contribution T_{sky} is quite low at L-band and set to a constant value of 3.7 K according to [21; 30].

To study the effect of surface roughness on the measured signal, the prevailing surface conditions are divided into four classes of differing σ . Ranges of σ are defined from a trend of measured σ (Eq. 5) to better emphasize the effect of roughness on the signal. The evolution of σ with time (Table I and Fig. 1) suggests the following ranges : $\sigma < 16$ mm relative to smooth surface, i.e. before the campaign ; σ belonging to the range 16-20 mm characterizing the steady state reached by the surface at the end of the campaign, from the end of April 2006 to March 2007 ; σ between 20 and 24 mm for the transition between very rough and steady state surface, from February to April 2006 ; and a last case concerning a very rough surface characterized by a σ higher than 24 mm, just after ploughing.

B. Surface modeling

This database is also used to retrieve and study the semi-empirical parameters in the L-MEB that account for the effect of a rough surface [3; 7]. The emission of a flat surface is obtained by computing its dielectric constant from soil conditions, i.e. texture, temperature and surface soil moisture. The model developed by Mironov et al. [31; 32] is used as it has been shown to be more relevant for our experiment site [23] than the Dobson's model [33; 34]. The reflectivity $\Gamma = 1-\epsilon$, is then derived using Fresnel's law for a flat soil. The surface emission, or reflectivity, must then be corrected to take into account a rough air-soil interface. This roughness contribution is estimated by the following semi-empirical approach [10; 17; 18]:

$$\Gamma_p(\theta) = [(1 - Qr) \cdot \Gamma_p^0(\theta) + Qr \cdot \Gamma_q^0(\theta)] \cdot e^{-Hr \cdot \cos^{N_p}(\theta)} \quad (2)$$

where Γ is the reflectivity with the subscripts p and $q = V$ or H for the Horizontal and Vertical polarizations; the index 0 stands for reflectivity of a flat surface computed from the Fresnel's law; θ being the incidence angle; Qr , Hr , N_p are the roughness parameters to be calibrated [10; 17]. Qr is a mixing factor that allows us to take into account the polarization mixing caused by the rough surface, N_p allows us to account for the incidence angle [35] and depends on the polarisation [18] and Hr is the effective roughness parameter.

292 A first attempt to relate these empirical
 293 parameters to surface roughness suggested that
 294 $Hr = (2k\sigma)^2$ [9]. Hr was also found to depend
 295 on soil moisture [17; 18; 14] but as it has not
 296 been confirmed [23], it is not considered in the
 297 present study. This dependence could be partially
 298 explained by a mismatch between sampling depth
 299 of soil moisture sensors and the actual depth
 300 of the surface emission layer in L-band [23].
 301 N_p ($p = H$ or V for horizontal and vertical
 302 polarization) was found to be different for the two
 303 polarizations and $N_H=1$ and $N_V=-1$ were found
 304 for our SMOSREX site [18]. Q_r is generally
 305 considered to be negligible [14; 15; 13; 18] at
 306 L-band but in reality a rough surface implies
 307 mixing in polarization [10; 26] that can only be
 308 simulated by setting $Q_r > 0$ [11].

309
 310 Parameter retrieval:
 311 4 parameters are unknown in Eq. 2, that are Q_r , Hr ,
 312 N_H and N_V . The retrieval is done in two steps. The
 313 first one is based on a relationship between N_H
 314 and N_V [18]. Indeed, both theory using Fresnel
 315 law and observations over a flat surface show that
 316 the reflectivity at H and V polarizations are related
 317 by the following approximate equation (see [18])
 318

$$\Gamma_H(\theta) = [\Gamma_V(\theta)]^{\cos \Delta N(\theta)} \quad (3)$$

319 For a smooth surface, ΔN (Eq. 3), i.e. the
 320 difference ($N_H - N_V$), was found to be equal to
 321 2 [18] which is not relevant for a rough surface
 322 [11]. $\Gamma_H(\theta)$ and $\Gamma_V(\theta)$ are extracted from our
 323 database (i.e. LEWIS measurements) for each day
 324 of the roughness campaign (see Table I, left hand
 325 column) allowing us to compute ΔN for rough
 326 conditions. The second step uses Eq. 2 from Lewis
 327 brightness temperatures, where $N_H - N_V$ are linked
 328 together as a results of the first step.

329 The retrieval consists of minimizing a cost function
 330 that computes the quadratic differences between
 331 measured emissivities (ϵ_{lewis} at incidence angles
 332 of $\theta = 20, 30, 40, 50^\circ$ and both polarizations)
 333 and simulated emissivities (ϵ_{model}). This sets the
 334 best values of parameters (Eq. 2) that fit the
 335 observations [3] [5]. The cost function to be
 336 minimized is:
 337

$$CF = \frac{\sum (\epsilon_{lewis} - \epsilon_{model})^2}{\delta(\epsilon_{lewis})^2} + \sum_i \frac{\sum (P_i^{init} - P_i)^2}{\delta(P_i^2)} \quad (4)$$

338 where ϵ_{lewis} at all angles and polarizations
 339 are used; $\delta(\epsilon_{lewis})$ being the error in emissivity
 340 measured by LEWIS instrument [21]; P_i are

the retrieved parameters (Q_r , Hr , and N_p), P_i^{init}
 the initial values of the retrieved parameters
 (respectively $Q_r^{init} = 0.1$, $Hr^{init} = 0.75$, $N_p^{init} = 1$);
 and $\delta(P_i)$ the standard deviation of the retrieved
 parameters ($\delta Q_r = 1$, $\delta Hr = 2$, $\delta N_H = 1$).

As Q_r was found to be = 0 [13], two cases are
 considered here: A) where $Q_r = 0$ and Hr , N_H and
 N_V are retrieved and B) all the 4 parameters Q_r ,
 Hr , N_H and N_V are retrieved.

IV. RESULTS AND DISCUSSION

This section presents the results obtained from
 the SMOSREX-2006 campaign. Firstly, roughness
 measurements are presented for 14 months and
 secondly, the emissivities measured by the LEWIS
 instrument are analyzed to better understand
 the effect of roughness on the L-band surface
 emission. Finally, this database is used to study the
 semi empirical model that accounts for roughness
 in L-MEB. The parameters of the semi-empirical
 model are derived and related to surface roughness
 conditions.

A. Measured roughness

Table I presents the means and standard
 deviations of σ and LC as well as the ratio σ/LC
 acquired during σ each day of the campaign. Mean
 values are obtained considering samples at both
 orientations, i.e. parallel and perpendicular to
 LEWIS plane of incidence. Before ploughing,
 the surface was almost flat characterized by
 $\sigma=4.73\pm 1.31$ mm and a correlation length $LC =$
 94.11 ± 38.81 mm. As a comparison, previous
 measurements of the SMOSREX site [18] reported
 $\sigma = 11.09$ mm in February 2004 and $\sigma = 9.12$ mm
 in April 2004, indicating a smooth surface. After
 ploughing, the surface was characterized by a
 standard deviation height σ of $34.58 \text{ mm} \pm 10.29$
 mm and a correlation length of 62.42 ± 26.68 mm.
 The auto-correlation functions (Eq. 1) suggest that
 the surface is closer to an exponential one than a
 gaussian one [26; 27].

The time variations of σ (top panel), the cor-
 relation length (2^{nd} panel from the top), the soil
 moisture (3^{rd} panel from the top) from the end of
 2005 to March 2007 and the emissivity monitored
 at an incidence angle of 40 at both polarizations
 (bottom panel) are given in Fig. 1. The effects
 of the soil ploughing can be clearly distinguished
 on January 13th (top panel) and is characterized

TABLE I
STANDARD DEVIATION OF HEIGHTS, σ , AND THE CORRELATION LENGTH, LC, FOR EACH DAY OF THE CAMPAIGN. σ AND LC ARE AVERAGED FROM EVERY SAMPLES ACQUIRED AT BOTH DIRECTIONS. THE RIGHT-HAND COLUMN IS THE RATIO σ /LC

date Year mm/dd/yy	Roughness Characteristics		
	Standard Deviation of surface height σ (mm)	correlation length LC (mm)	σ /LC (mm)
02-07-03*	11.51* \pm 2.72	59.56* \pm 35.90	0.19*
02-04-04*	11.09* \pm 3.59	101.22* \pm 42.20	0.11*
04-02-04*	9.12* \pm 2.18	70.67* \pm 33.70	0.13*
01-13-06*	4.73* \pm 1.31	94.11* \pm 38.81	0.05*
01-13-06	34.58 \pm 10.29	62.42 \pm 26.68	0.55
01-20-06	29.67 \pm 9.66	70.21 \pm 29.55	0.42
02-01-06	26.85 \pm 11.17	60.99 \pm 16.90	0.44
02-20-06	25.58 \pm 5.86	65.26 \pm 22.88	0.39
03-16-06	23.10 \pm 6.61	76.06 \pm 33.78	0.30
04-03-06	25.44 \pm 6.76	87.78 \pm 34.97	0.29
05-04-06	20.93 \pm 7.05	96.08 \pm 56.66	0.22
05-30-06	20.32 \pm 7.22	82.39 \pm 31.60	0.25
06-29-06	18.05 \pm 4.84	105.19 \pm 43.16	0.17
11-24-06	19.25 \pm 5.99	118.21 \pm 33.12	0.16
03-12-07	17.43 \pm 5.72	115.32 \pm 42.66	0.15
10-06-10	12.31 \pm 3.19	122.68 \pm 62.42	0.10

* Measurements before ploughing

393 by a sharp increase in σ followed by a noticea24
394 ble decrease in σ from January to May. Then σ 5
395 decreases more slowly, reaching a quasi-constant6
396 value by July 2006. After 14 months σ was about7
397 17.4 mm. In June 2010 the soil roughness was mea28
398 sured (Table I) and presented a level of roughness9
399 comparable with the value measured in April 2004,0
400 as $\sigma=12.31\pm 3.19$ mm and $LC=122.68\pm 62.42$ mm,1
401 This trend is well reproduced using an exponential,2
402 fit function (dashed line top panel Fig. 1) as: 433

$$\sigma = 38.35xDOE^{-0.126} \quad (434)$$

403 with DOE being the Day of the Experiment (dashed436
404 line top panel Fig. 1). The correlation length 437
405 LC- presents an opposite behaviour, showing a low438
406 value after ploughing and increasing with time as439
407 the surface becomes less and less rough. A fit440
408 function was used to represent its trend (dashed441
409 line, 2nd panel from top Fig. 1) and is defined as442

$$LC = 48.67xDOE^{0.132} \quad (443)$$

410 The effect of ploughing leads to a decrease in444
411 soil moisture as shown in Fig. 1 (2nd Fig. from445
412 the bottom) in January 2006. This effect could be446
413 explained by a redistribution of the water content447
414 within the soil. Consequently, the emissivity448
415 (bottom panel of Fig. 1) increases whereas the449
416 difference of polarization, $\epsilon_V - \epsilon_H$, decreases. It450
417 should be noted that ploughing changes also the451
418 bulk density: the soil density decreasing from452
419 1.5 kg/m³ in 2005, to 1.39 kg/m³ in February453
420 20th, 2006. Weather conditions then compact the454
421 surface, decreasing σ and increasing the density to455
422 1.57 kg/m³ in November 2006. Thus, ploughing456
423 the surface modifies the soil properties (bulk457

density, soil moisture redistribution) impacting the
dielectric constant and so the surface emissivity
[17].

σ and LC are correlated as seen in Fig. 2, which
reports the relation existing between LC, σ /LC and
 σ^2 /LC as a function of σ . Estimating LC from field
measurements is difficult (i.e. the measurements
are noisy) but a modeling study [36] has shown
that it has a very low influence on brightness
temperature, especially at H polarization. The
results of σ and LC are slightly different to what
was obtained with the same database [26] as their
methodology to compute σ and LC is different.

B. Observations of surface emissivities

Fig. 3 presents the emissivity calculated from
LEWIS measurements as a function of soil moisture
at 4 incidence angles, from $\theta=20^\circ$ (top row)
to $\theta=50^\circ$ (bottom row) and for both polarizations
(V black dots and H grey dots). The different
columns correspond to the four roughness classes
from quasi-smooth on the right to rough surfaces on
the left. Emissivity computed from Fresnel's law is
plotted (grey and black lines Fig. 3) characterizing
the emission of a perfectly smooth surface with
identical surface conditions (i.e. with the same
soil moisture, density, temperatures). As expected,
emissivity decreases with increasing soil moisture
at both polarizations and all angles. The effect of
roughness is to decrease the sensitivity of surface
emission to soil moisture. This can be observed
especially at wet conditions (i.e. $> 0.25m^3/m^3$),
where the emissivity increases with roughness. The

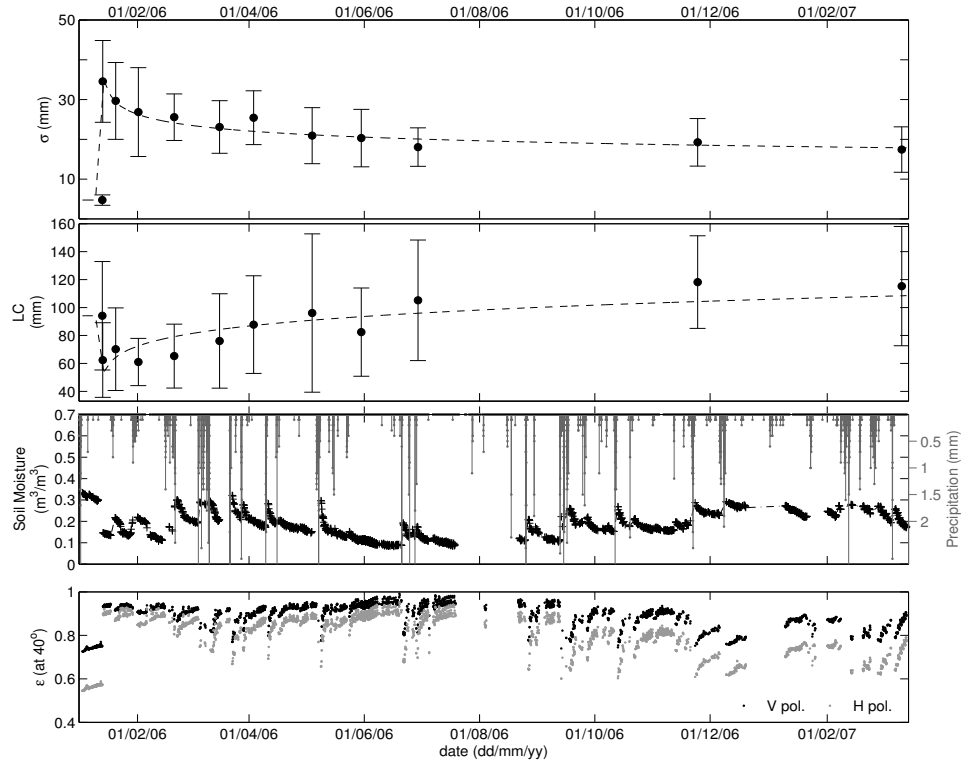


Fig. 1. Time series of surface parameters from December 2005 to March 2007. Top Fig. is σ (in mm) and its standard deviation ; The surface was ploughed the 13th of January 2006. 2nd from the top: the correlation length ; 3rd panel: soil moisture (black x, left hand y-axis) and precipitation (grey sticks, right hand y-axis, note that it is inverted for graphical convenience) ; Bottom figure is the emissivities at V (black dots) and H (grey dots) polarizations monitored by Lewis radiometer at an incidence angle of 40° .

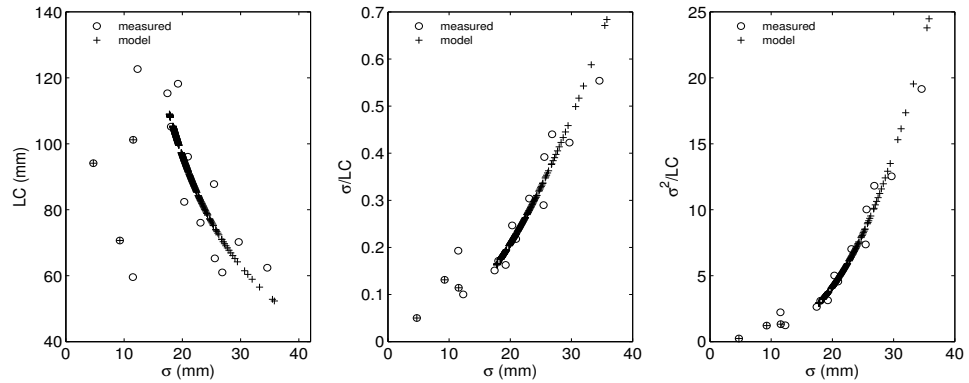


Fig. 2. LC , σ/LC and σ^2/LC as a function of σ . For each case are displayed: the measured σ and LC (Table I), “o” symbols and referred to as “measured” in the legends; σ and LC obtained from Eq. 5 and 6, “+” symbols and referred to as “modeled” in the legends. Measured and model data are similar for data acquired before ploughing the surface.

458 difference between the emissivities at H and V pol-
 459 arization increases with increasing incidence angle
 460 for each wetness conditions but is decreased with
 461 roughness. Furthermore, the impact of roughness
 462 on the emissivity is more pronounced at H than
 463 V polarization. At the incidence angle of 40° , the
 464 emissivity at H pol. is ~ 0.56 at a soil moisture
 465 content of $0.3m^3/m^3$ and for a smooth surface (3rd
 466 line, right hand side Fig. 3) whereas it is ~ 0.8 for a
 467 rough surface (left hand side Fig. 3). It corresponds

to an increase in the emissivity of 0.24, whereas for
 the V polarization this increase is ~ 0.145 , from an
 emissivity of ~ 0.72 for flat condition to ~ 0.865 for
 a rough surface. The decrease in the emissivity with
 soil moisture has a linear trend for rough conditions
 and for each incidence angle (left-hand columns
 Fig. 3), the effect being again more pronounced at
 H polarization than at V polarization.

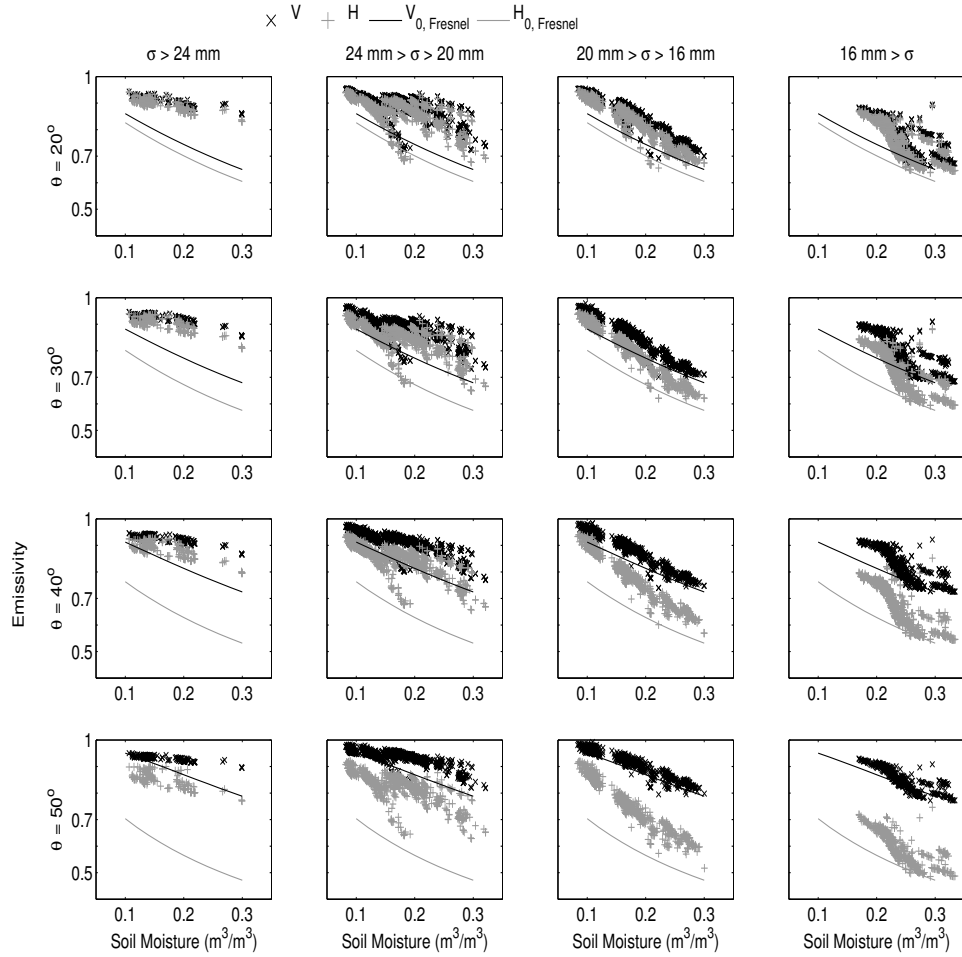


Fig. 3. Emissivity at V (black x) and H (grey +) polarizations, monitored at 4 incidence angles as a function of soil moisture : 20° (top row figures), 30° (2nd row), 40° (3rd row) and 50° (bottom row). The 4 columns correspond to roughness conditions, from a very rough surface -1st column from the left- to quasi-smooth condition (right hand side column). Emissivity computed from Fresnel's law (flat surface) is shown as black (V pol.) and grey (H pol.) continuous lines.

476 C. L-Meb model calibration

477 The second objective of this paper is to use the
478 database to study the roughness parameters (Q_r ,
479 H_r , N_H and N_V) as defined in Eq. 2.

480
481 1) Relation between N_H and N_V : ΔN is derived
482 from Eq. 3 and presented in Fig. 4 as a function of
483 σ values estimated by the fit function (Eq. 5 and
484 grey dashed line Fig. 1). The use of the fit instead
485 of actual values is done to limit errors caused by
486 sampling limits in characterizing the field (2196
487 board and ~ 8 samples per day). Fig. 4 clearly
488 shows a decreasing trend of ΔN with σ , well
489 represented by the linear function defined as ΔN
490 $= N_H - N_V = -0.049 \cdot \sigma + 2.188$ ($R = 0.90$, $RMSE$
491 $= 0.16$, $bias=0$). Smoother surface, i.e. $\sigma < 16$
492 mm, is characterized by a ΔN of ~ 1.8 , which is
493 in agreement with $\Delta N = 2$ found previously [18]
494 whereas it is ~ 0.5 for very rough surface, i.e. $\sigma > 4$
495 35mm. This trend is close to that obtained in [13]

496 ($\Delta N = -0.036 \cdot \sigma + 2.24$) over another agricultural
site.

2) Retrieved parameters: N_p ($p= H$ or V), Q_r
and H_r (Eq. 2) were derived from Eq. 4, for every
day over the period November 2005-April 2007.
The emissivity computed using these parameters,
leads to an $RMSE=0.022$ ($R^2=0.95$) when compared
to LEWIS emissivity, whereas an $RMSE = 0.053$
($R^2=0.69$) is encountered when applying the
parameters found by Escorihuela et al. [18] over a
flat surface. Fig. 5 presents the retrieved roughness
parameters Q_r (top Fig.), H_r (middle Fig.) and
 N_H and N_V (bottom Fig.) for case B as a function
of time. The time variation in σ and its best fit
trend (Eq. 5) are also showed for comparison. H_r
presents a high variability, but in general it decreases
as σ decreases.

The high variability in the retrieved values of
 H_r could be linked to the fact that this parameter

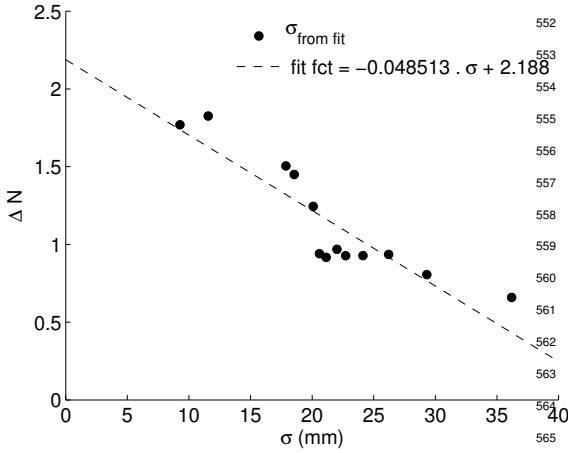


Fig. 4. σ estimated from roughness measurements plotted against ΔN (black \bullet) calculated from LEWIS data, including the linear fit function (dashed line).

516 tends to compensate for the difference between
 517 the sampling depth [23] [37] of the in-situ soil
 518 moisture sensors ($\sim 0-6$ cm top soil layer),
 519 and of the LEWIS observations ($\sim 0-2/3$ cm).
 520 For example after a rain event following dry
 521 soil moisture conditions, the LEWIS observations
 522 immediately show a clear decrease in the monitored
 523 brightness temperatures whereas the in-situ probe
 524 still measures a low water content. Whilst LEWIS
 525 is sensitive to the first 0-2/3 cm, which is wet after
 526 a rain event, the probe integrates the soil moisture
 527 between the surface layer which is wet and a
 528 deeper layer which is dryer. In this case, the soil
 529 moisture estimated by the probe is underestimated
 530 in comparison to the soil moisture seen by LEWIS.
 531 The L-MEB model uses this underestimated soil
 532 moisture and compensates this effect by adjusting
 533 H_r to fit the LEWIS observations. Such effects
 534 may explain the high variability in the retrieved
 535 values of H_r obtained in May, July, September
 536 2006. The opposite situation is also observed
 537 (dry surface over the 0-2/3 cm surface layer and
 538 rather wet conditions over the 0-6 cm surface
 539 layer) and could explain high retrieved values
 540 of H_r obtained in March 2004 and November 2005.
 541
 542 The results of the retrieval are presented as a
 543 function of the estimated σ (Eq. 5, dashed line
 544 top Fig. 1) and LC (Eq. 6) in Fig. 6 and Fig. 7
 545 (grey markers for the case A with $Q_r=0$ and black
 546 markers for the case B where Q_r is retrieved).
 547 We also studied the derived parameters with the
 548 quantity σ/LC (not shown here), but the results
 549 are very similar to the results presented in Fig. 6.
 550 Q_r (case B, it is retrieved, black \bullet Top left Fig. 6)
 551 increases significantly from values around 0.05 for

a flat surface to 0.3 for a rough surface. A Low Q_r
 value for a quasi-smooth surface is in agreement
 with both theory (no polarization mixing, [11])
 and observations [13] [18]. It confirms also that
 Q_r is not equal to 0 for rough surface and needs
 to be taken into account to model the signature
 of rough soils. Retrieved values of H_r (Top right
 Fig. 6) show more variability as mentioned earlier.
 They evolve on average from $\sim 0.2-0.3$ for a
 smooth surface to ~ 1 for a rough surface. The
 relation $H_r=f(\sigma)$ obtained in [13] is represented by
 the dashed line, fitting the results of the presented
 study. It is interesting to note that this relationship
 obtained for different conditions over a different
 site and a variety of soil roughness conditions
 provide a good general fit to the results obtained in
 this study. These results confirm that the empirical
 relationship $H_r = (2k\sigma)^2$ [9] (dotted line Fig. 6)
 is not applicable, also found in [13]. Retrieved
 values of H_r when Q_r , H_r and N_p ($p = H$ or V)
 are retrieved are higher than when Q_r is set equal
 to 0. Q_r and H_r variations seem to be correlated to
 variations in σ whereas no clear correlation with
 σ could be found for N_V and N_H (bottom left
 Fig. 6) confirming the observations of [13]. N_H
 and N_V are found on average to be equal to 2.8
 and 1 respectively for a smooth surface whereas
 the authors of [18] set them to lower values of 1
 and -1. For rough surface however, N_H and N_V
 do not vary and can clearly be set to $N_H= 0.59$ and
 $N_V=-0.30$. Q seems related to H_r (bottom right
 Fig. 6) by the relation $H=2.69*Q$ ($R=0.71$). Eq. 2
 imposes the conditions $Q=0$ for $H=0$, meaning the
 emissivity of a flat surface is that from Fresnel's
 law.

The retrieved parameters show the opposite
 behavior when studied as a function of LC (Fig. 7)
 with H_r and Q decreasing with increasing LC. N_V
 and N_H present less variations for a rough surface
 (low LC) than in Fig. 6.

V. CONCLUSIONS AND PERSPECTIVES

Roughness effects at L-band are complex and
 need more investigations to be fully understood
 and modeled [13; 38]. This paper presents the
 unique SMOSREX-2006 experimental database
 dedicated to study the effect of roughness at
 L-band over 14 months. A bare soil has been
 significantly ploughed at the SMOSREX site
 and continuously monitored by LEWIS L-band
 radiometer. It has been found that the influence
 of roughness is more important at high incidence
 angles (about 40 to 50°), high soil moisture values
 and at H polarization.

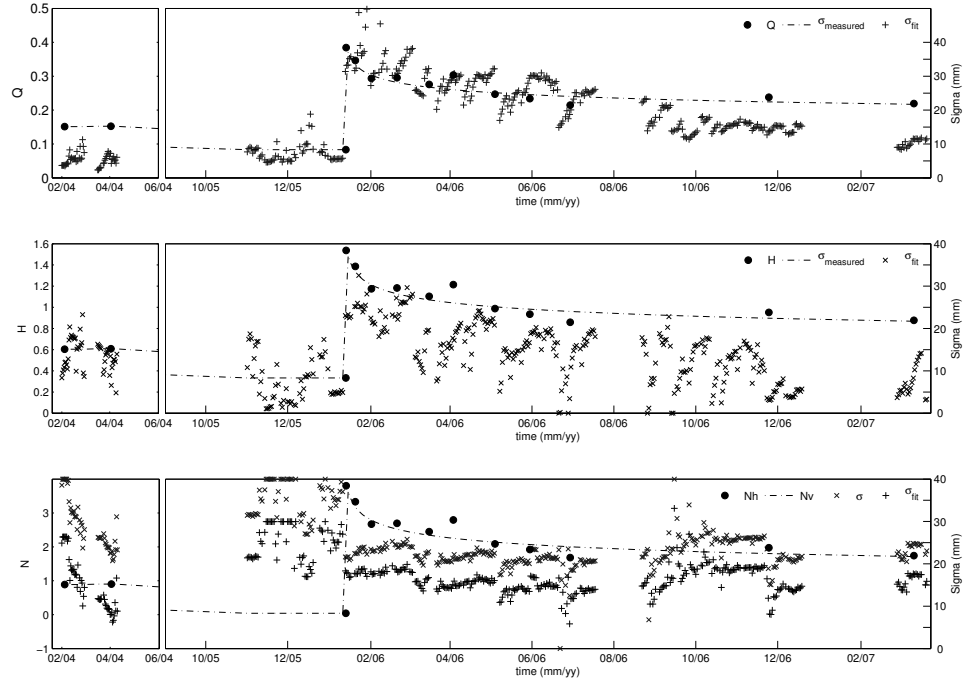


Fig. 5. Time series of the roughness parameters of Eq. 2. Top figure shows Q_r , middle figure shows H_r and the bottom figure presents N_H (x) and N_V (+). σ and its fit trend (Eq. 5) are also depicted (with right hand side y-axis). The time series are split in two panels (left and right hand columns) as the time series are not continuous (no roughness measurement in 2005).

607 The soil moisture derived from the SMOS₆
 608 mission is based on a semi-empirical approach₇
 609 [8] and the roughness effect is taken into account₈
 610 by the Q-H model [9; 13; 18]. The presented₉
 611 database is also used to study the semi-empirical₀
 612 parameters of the L-MEB emission model as a₁
 613 function of surface characteristics represented₂
 614 by σ and LC. The results of this study suggest₃
 615 that for a rough surface $Q_r=0.3$, $H_r \sim 1$, $N_H \sim$
 616 0.59 and $N_V = -0.30$, whereas a smooth surface₅
 617 is characterized by $Q_r \sim 0.05$, $H_r \sim 0.2/0.3$, N_H
 618 ~ 2.8 and $N_V \sim 1$. It is different from most of
 619 the previous works on the subject which set $Q=0$
 620 even for rough conditions. A simple model can
 621 not have been found to represent the dependence₈
 622 of the semi-empirical parameters with σ and LC₉
 623 due to their high variability, especially in case₀
 624 of H_r . However, it is interesting to note that₁
 625 the σ - H_r relation proposed by [13] seems to be₂
 626 applicable here over SMOSREX conditions. A
 627 linear relationship between N_H and N_V is also₃
 628 found, with the difference $N_H - N_V$ decreasing₄
 629 with σ . The variations of these semi-empirical₅
 630 parameters can be explained by the difference₆
 631 in sampling depth between the sensors that are₇
 632 not sensitive to the same surface layer. This₈
 633 difference can be reduced by selecting some₉
 634 certain weather and soil moisture conditions.
 635 After an important rainfall the soil reaches its₀

field capacity and is more homogeneous in terms
 of soil moisture content as both the 0-2/3 cm
 top layer (as monitored by LEWIS) and the top
 0-6cm (as monitored by the probes) should have
 the same soil moisture content. After a drying
 period, the soil reaches its lower soil moisture
 content and both the probes and LEWIS monitor
 the same amount of soil moisture. By extracting
 those specific periods, it is expected to reduce the
 variability of the derived parameters.

ACKNOWLEDGMENT

The authors would like to thank the SMOSREX
 partners Météo-France and ONERA. This ex-
 periment was funded by Programme National de
 Télédétection Spatiale and Terre Océan Surfaces
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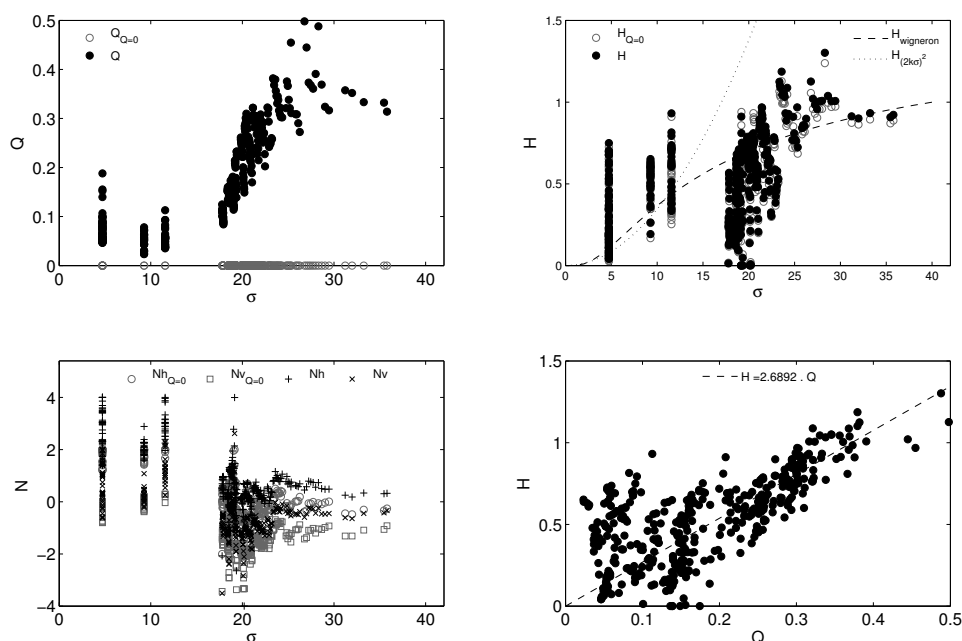


Fig. 6. Retrieved Q_r , H_r and N_p ($p=V$ or H) as a function of estimated σ (i.e. Eq. 5). Two cases considered : A) $Q_r=0$, H_r and N_p ($p=V$ or H) are derived (grey markers on all Figure); B) Q_r is derived (black markers). In the top right figure are also depicted H functions as found in i) Wigneron et al. 2011 [13] (dashed line) and ii) $H_r = (2k\sigma)^2$ [9] (dotted line).

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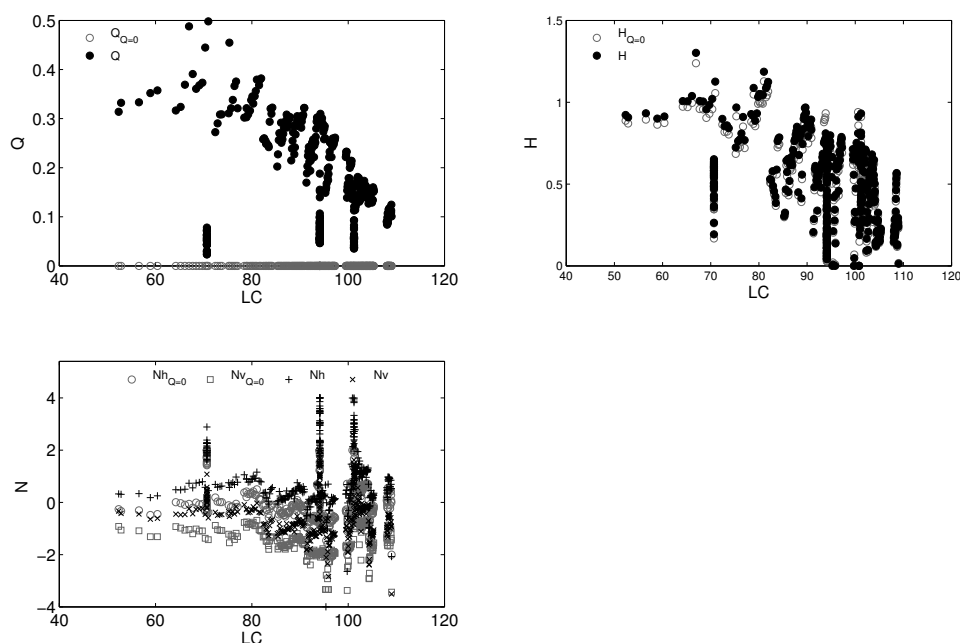
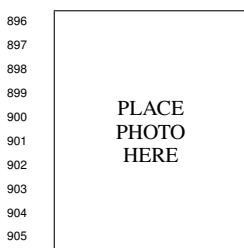


Fig. 7. Retrieved Q_r , H_r and N_p ($p=V$ or H) as a function of estimated LC (i.e. Eq. 6). Two cases considered : A) $Q_r=0$, H_r and N_p ($p=V$ or H) are derived (grey markers) B) Q_r is derived (black markers on all Fig.).

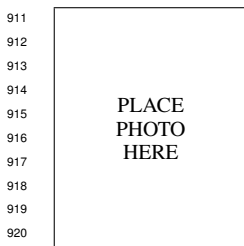
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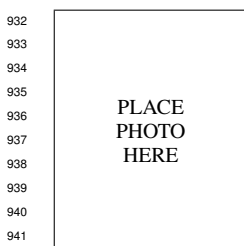
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895 press.



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941 fields of interest are in the theory and
942 techniques for microwave and thermal
943 infrared remote sensing of the Earth, with emphasis on hydrology,
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946 He was the science lead on the MIRAS project for ESA, and is now the
947 of the SMOS mission Lead-Investigator.



Patricia de Rosnay Patricia de Rosnay received her Ph.D degree in climate modelling from the University Pierre et Marie Curie (Paris 6, France) in 1999. She is currently Senior Scientist with the European Centre for Medium-Range Weather Forecasts (ECMWF) where is responsible of the land surface data assimilation activities. Her current research interests focus on improving soil moisture and snow analysis for Numerical Weather Prediction. She developed and implemented a 2-Dimensional Optimum Interpolation snow analysis and also implemented in operations an Extended Kalman Filter soil moisture analysis at ECMWF and she is investigating the use of passive and active microwave satellite data for Numerical Weather Prediction. She is involved in the SMOS Validation and Retrieval Team and participates to the EUMETSAT H-SAF project. She is member of the Science Definition Team and the Application Working Group of the future NASA SMAP mission and member of the SRNWP surface team. She was involved in land surface modelling activities such as the African Monsoon Multidisciplinary Analysis (AMMA) Land Surface Model Inter-comparison Project (ALMIP) and she coordinated the microwave component of the project ALMIP-MEM (Microwave Emission Model). She initiated the project for the validation of the future SMOS soil moisture products over West Africa. Patricia de Rosnay worked five years with the French Centre National de la Recherche Scientifiques (CNRS) at Centre d'Etudes Spatiales de la Biosphère (CESBIO), Toulouse, France, where she coordinated the SMOSREX field experiment in preparation of SMOS. Her research topic at the Laboratoire de Météorologie Dynamique (LMD) from 1994 to 2002 was focused on developments of global scale land surface processes representation in climate models.



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